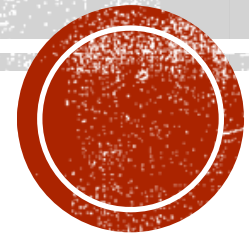


Workshop on Summer Low Flows in Western Oregon:
Processes, Trends, Uncertainties, and Forest Management

EVAPOTRANSPIRATION VARIABILITY ACROSS DOMINANT ECOSYSTEMS IN THE PNW

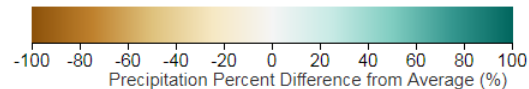
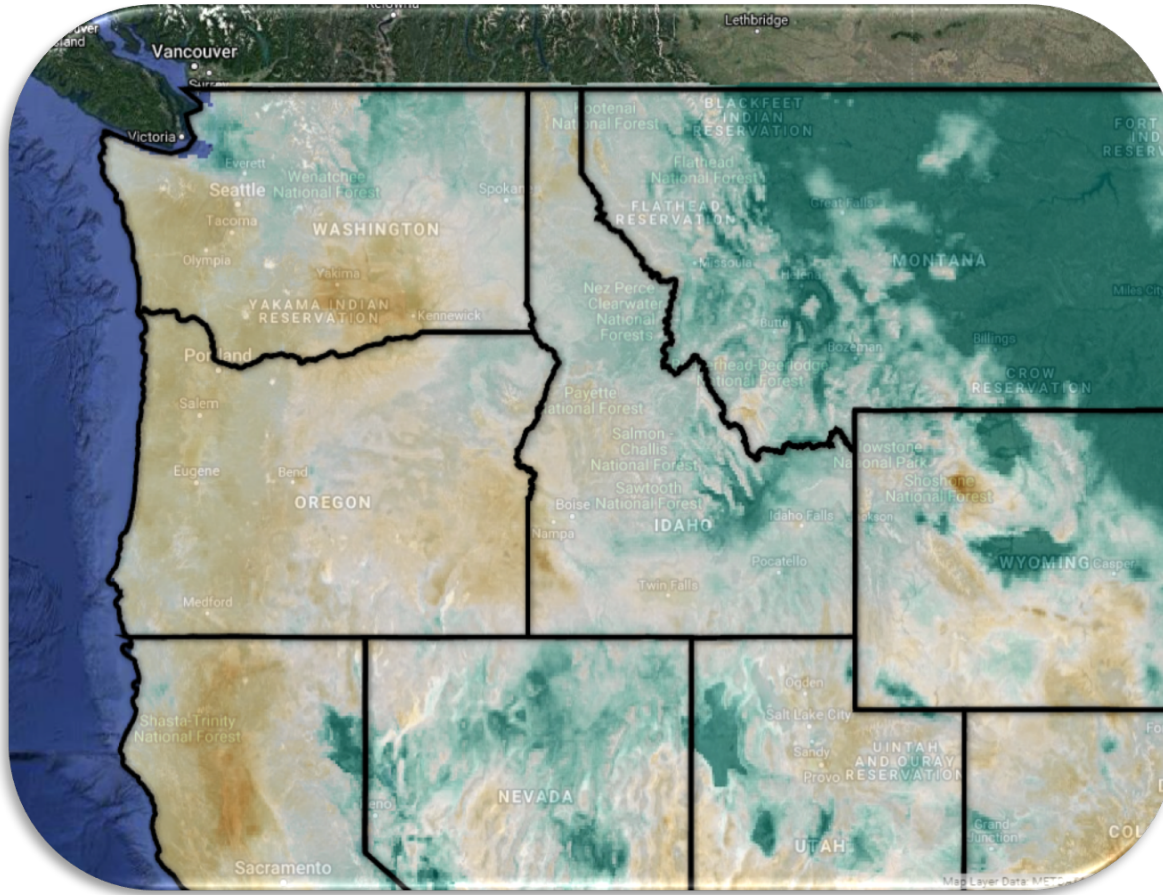
Hyojung Kwon, Beverly E. Law, and Christoph K. Thomas

Oregon State University



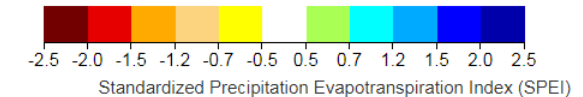
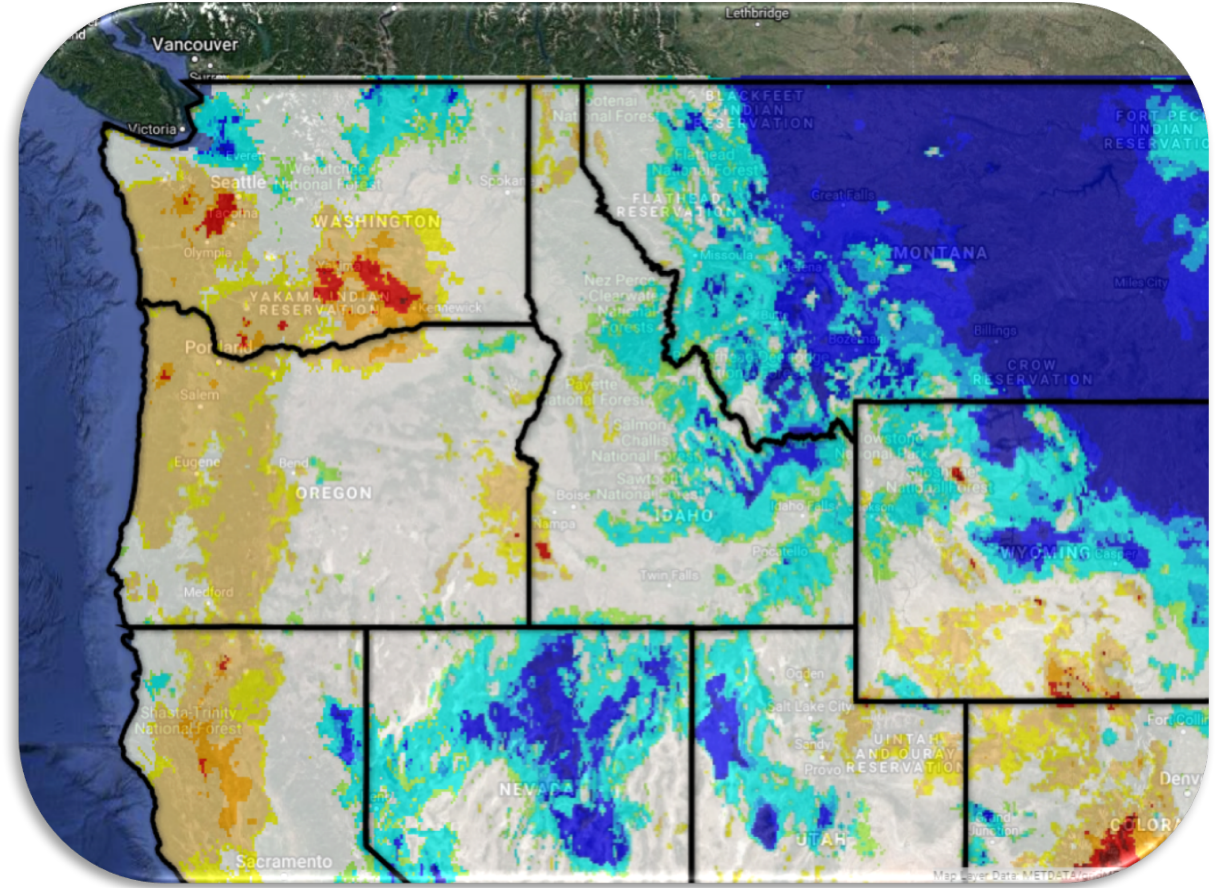
How Does Precipitation Look So Far This Year?

Precipitation



Drought Index (SPEI)

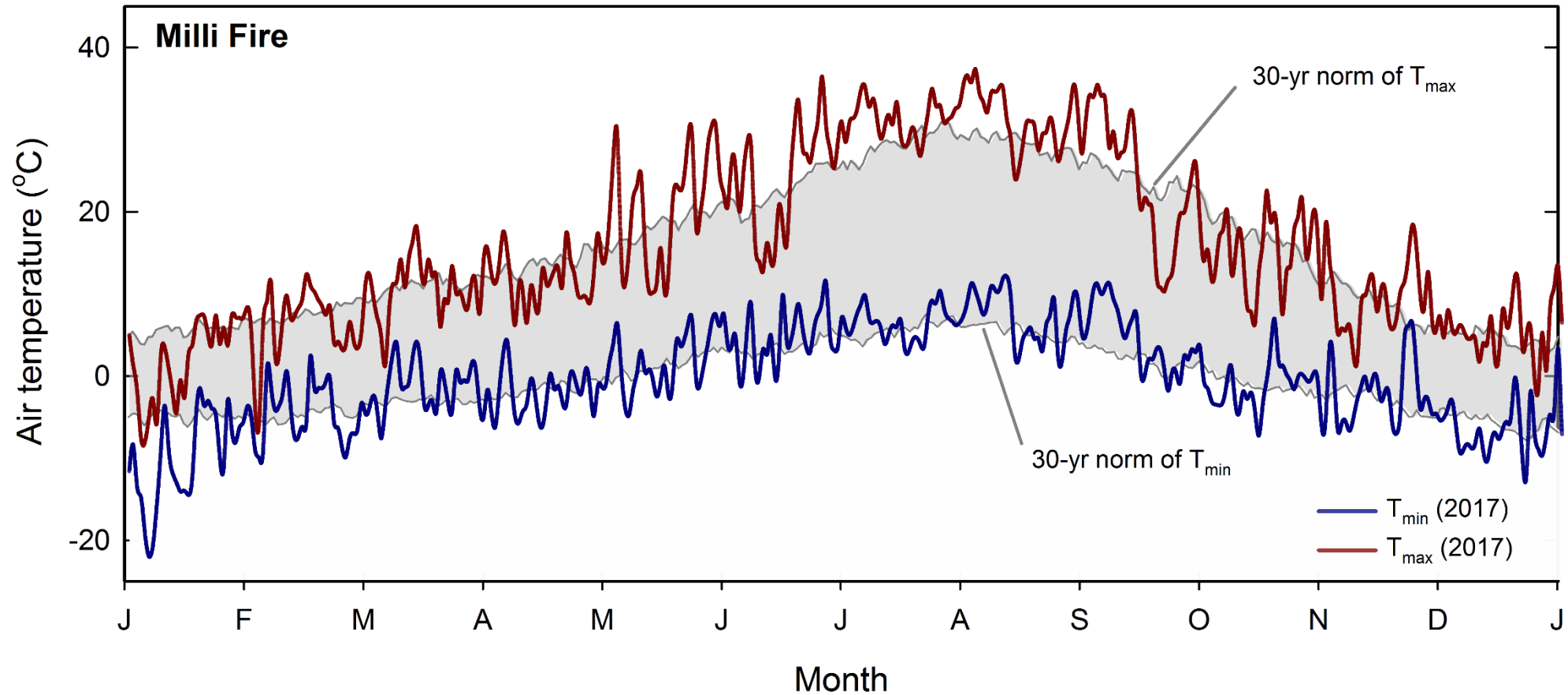
Standardized Precipitation Evapotranspiration Index



source: <https://app.climateengine.org/>

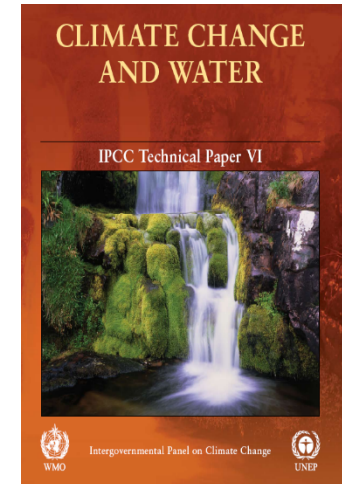


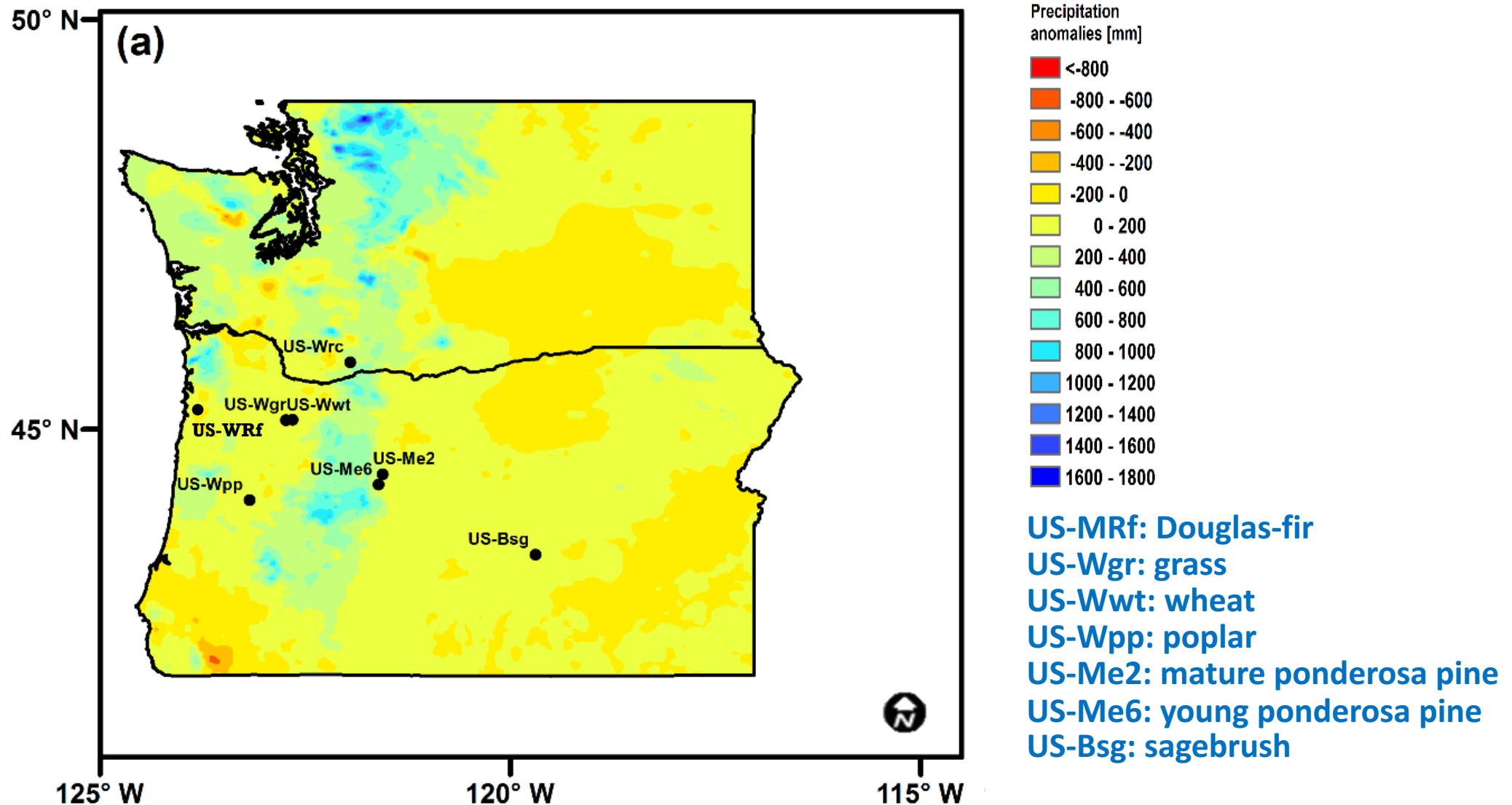
Increased T_{\max} And T_{\min}



*Progress in knowledge depends on **improved data availability** ... Relatively short records may not reveal the full extent of natural variability and confound detection studies ... **Major gaps in observations** of climate change related to freshwater and hydrological cycles are:*

precipitation, streamflow, soil moisture, and
actual evapotranspiration





Strong gradient from the coast to the inland



Scientific Objectives

- Characterize and quantify evapotranspiration of different ecosystems (from high desert sagebrush steppe to ponderosa pine to Douglas-fir) across temperature and hydrological gradients
- Assess the mechanism (climatological and biological components) that controls evapotranspiration



Focus of the presentation are:

- on seasonal patterns and controlling mechanism of evapotranspiration at three forest sites
- on evapotranspiration variation at seven sites

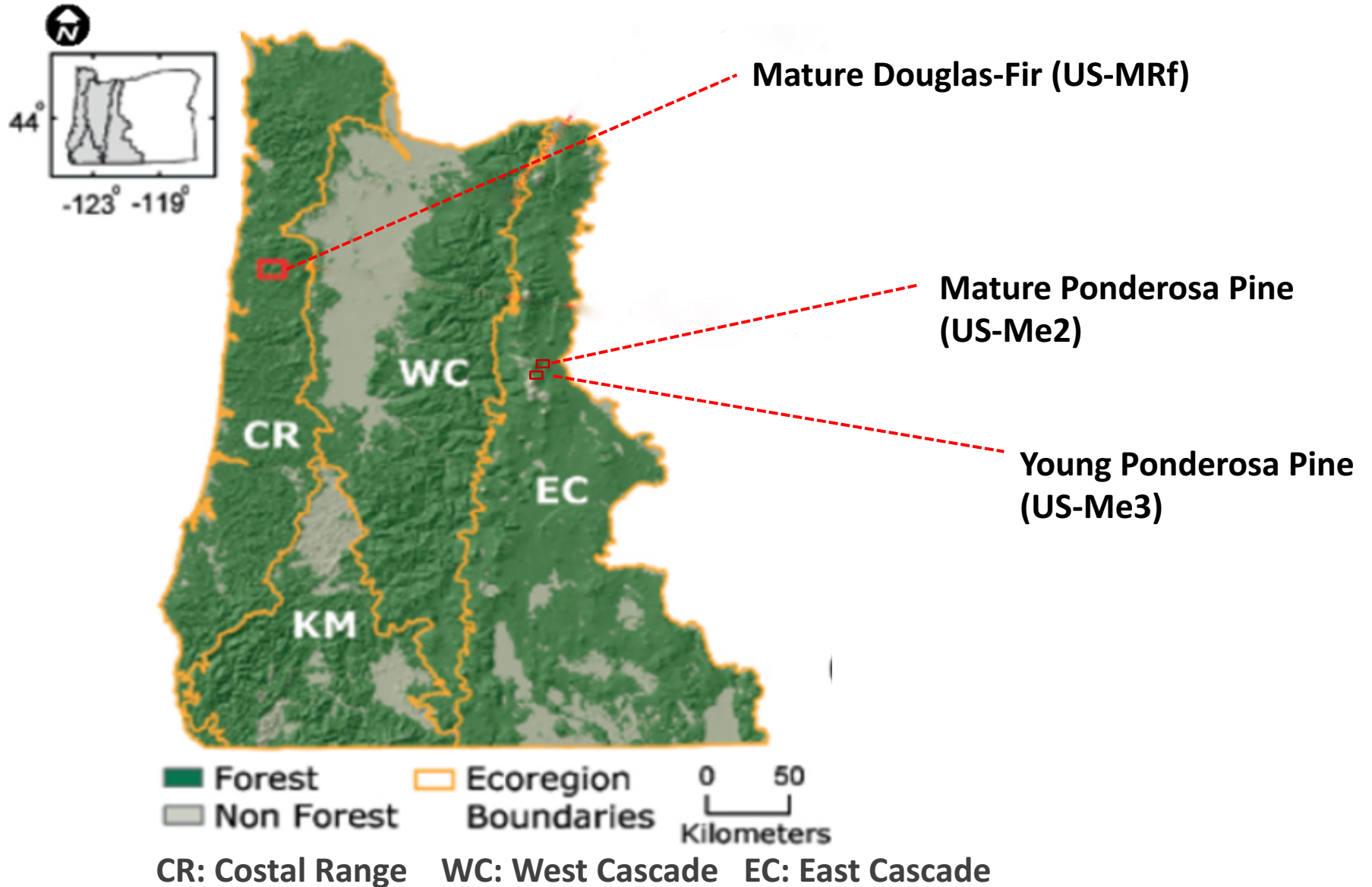


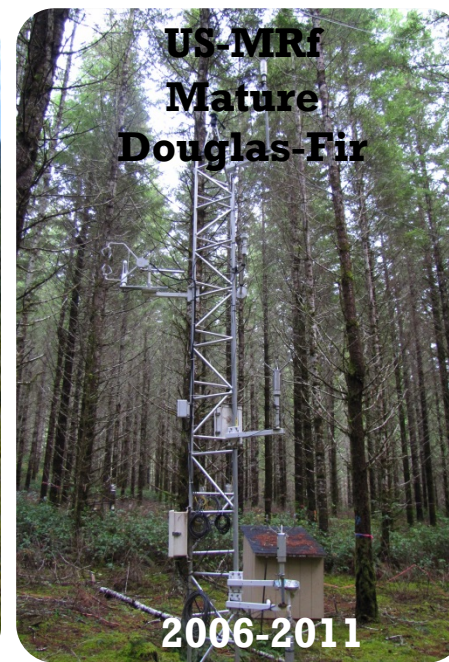
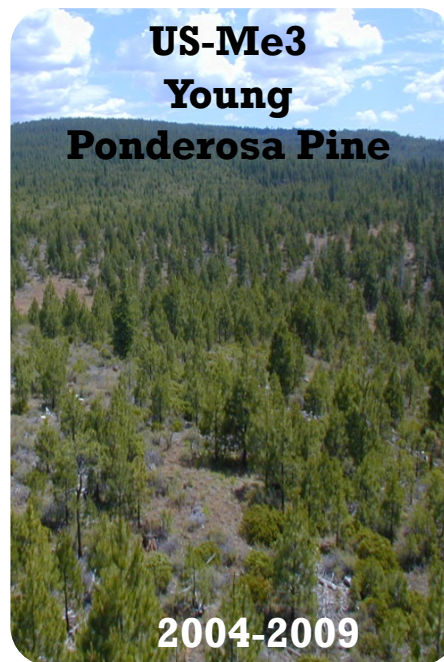
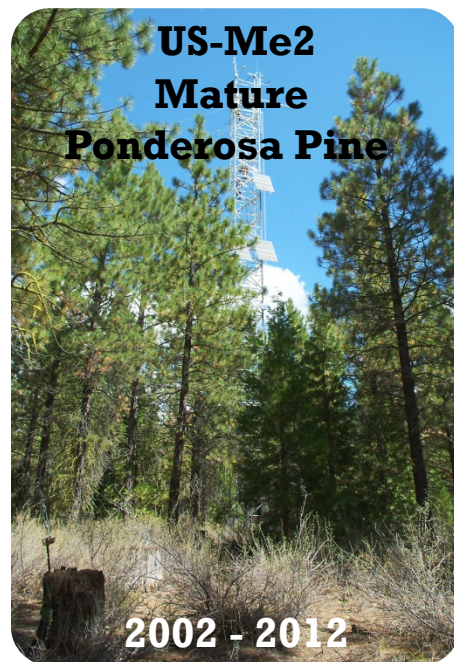
Evapotranspiration = **E**vaporation + **T**ranspiration

Precipitation = **ET** + runoff + ground water + storage
Net radiation = **λE** + sensible heat + ground heat + storage



Study Sites





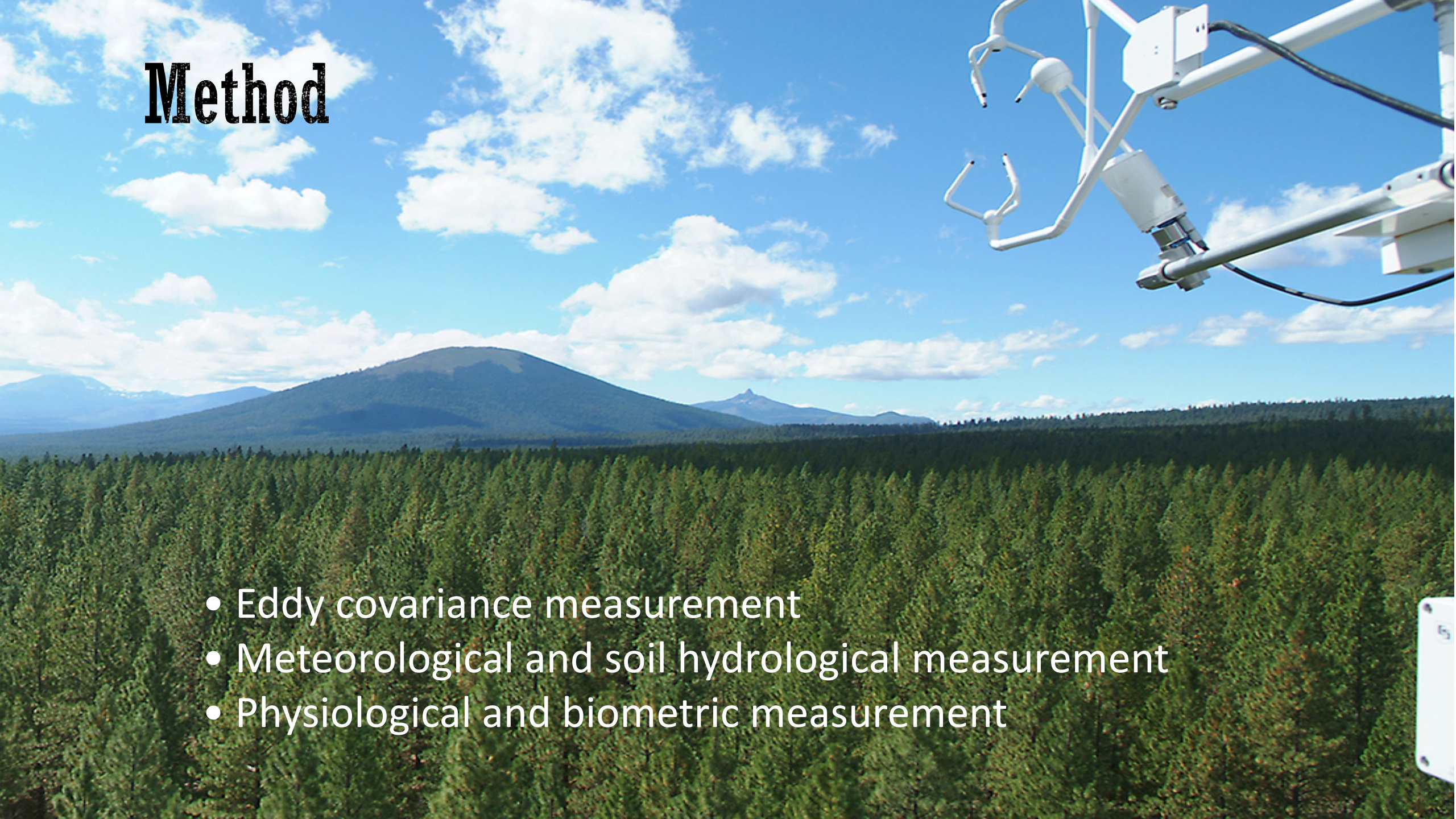
	US-Me2	US-Me3	US-MRf
Mean age (years)	64	16	58
Mean height (m)	22.0	3.3	34.0
Density (trees ha ⁻¹)	325	260	400
Maximum LAI	2.8	1.8	9.0
30-yr norm T _{air} (°C)	7.5	7.6	10.6
30-yr norm Prec (mm)	530	494	1820

Source of 30-yr normal: PRISM (Parameter-elevation Regressions on Independent Slopes Model) data

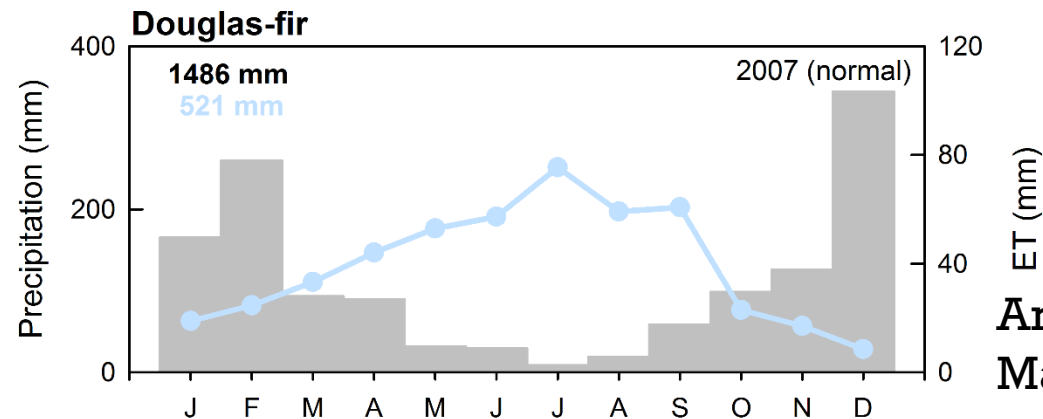
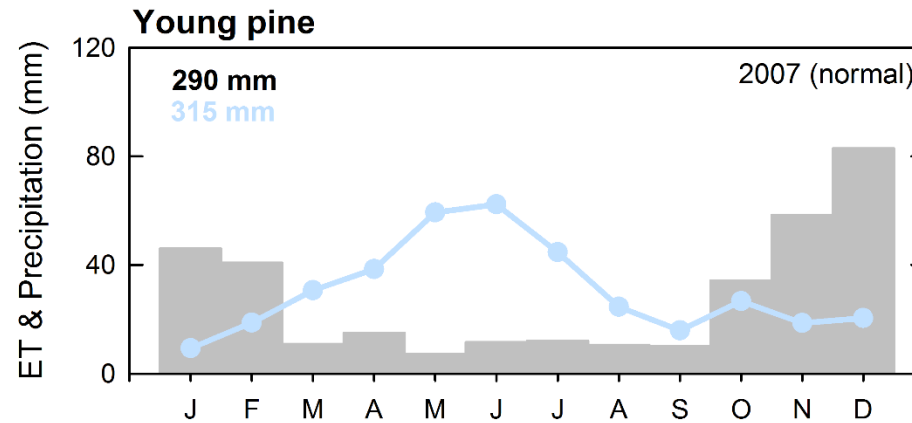
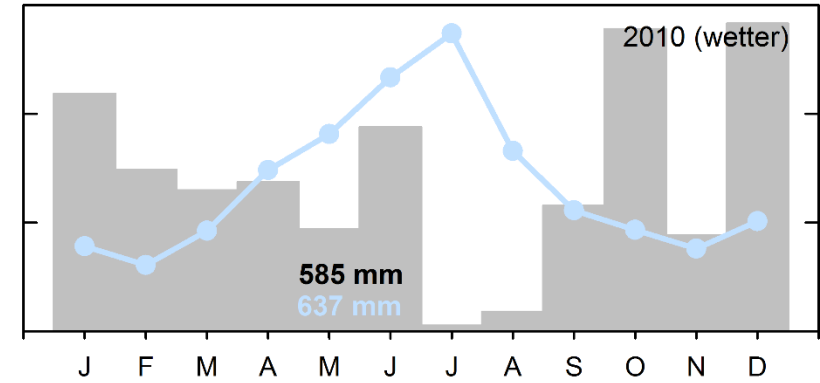
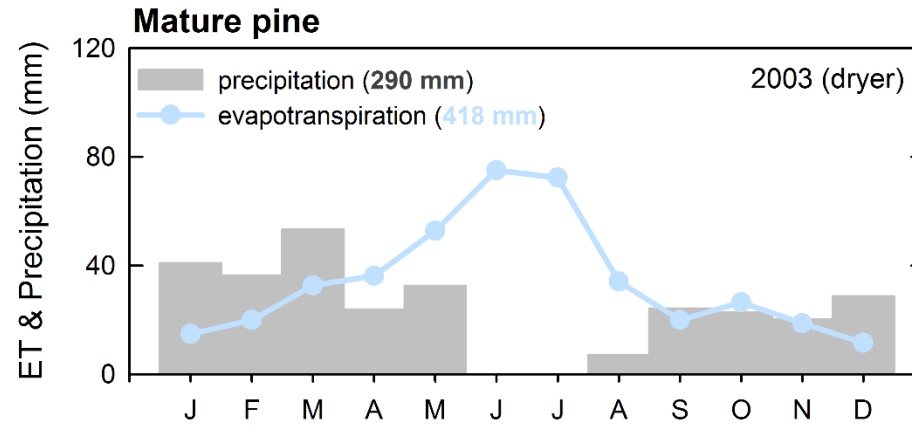


Method

- Eddy covariance measurement
- Meteorological and soil hydrological measurement
- Physiological and biometric measurement



Seasonal Pattern

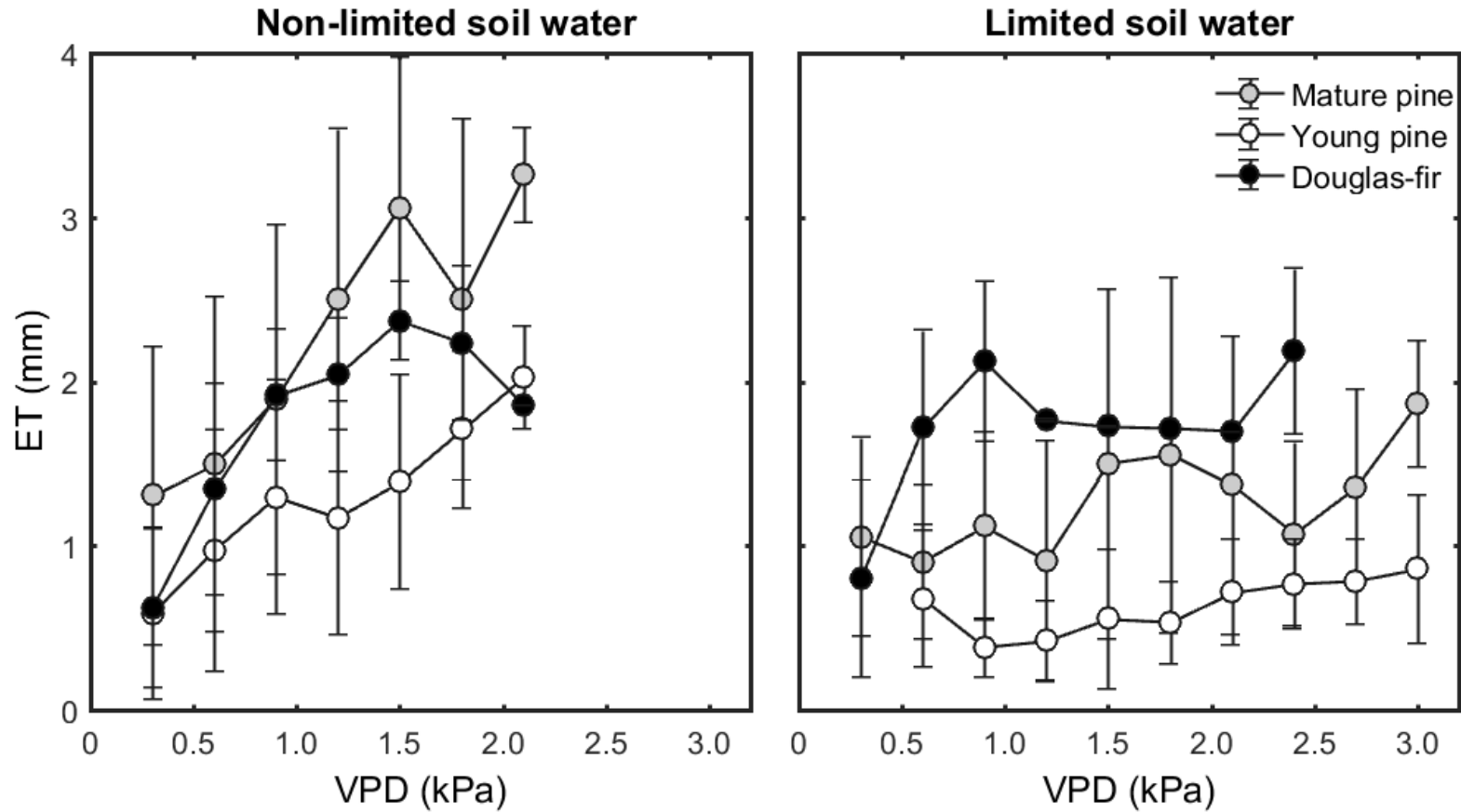


Annual ET:

Mature pine > Douglas-fir > Young pine



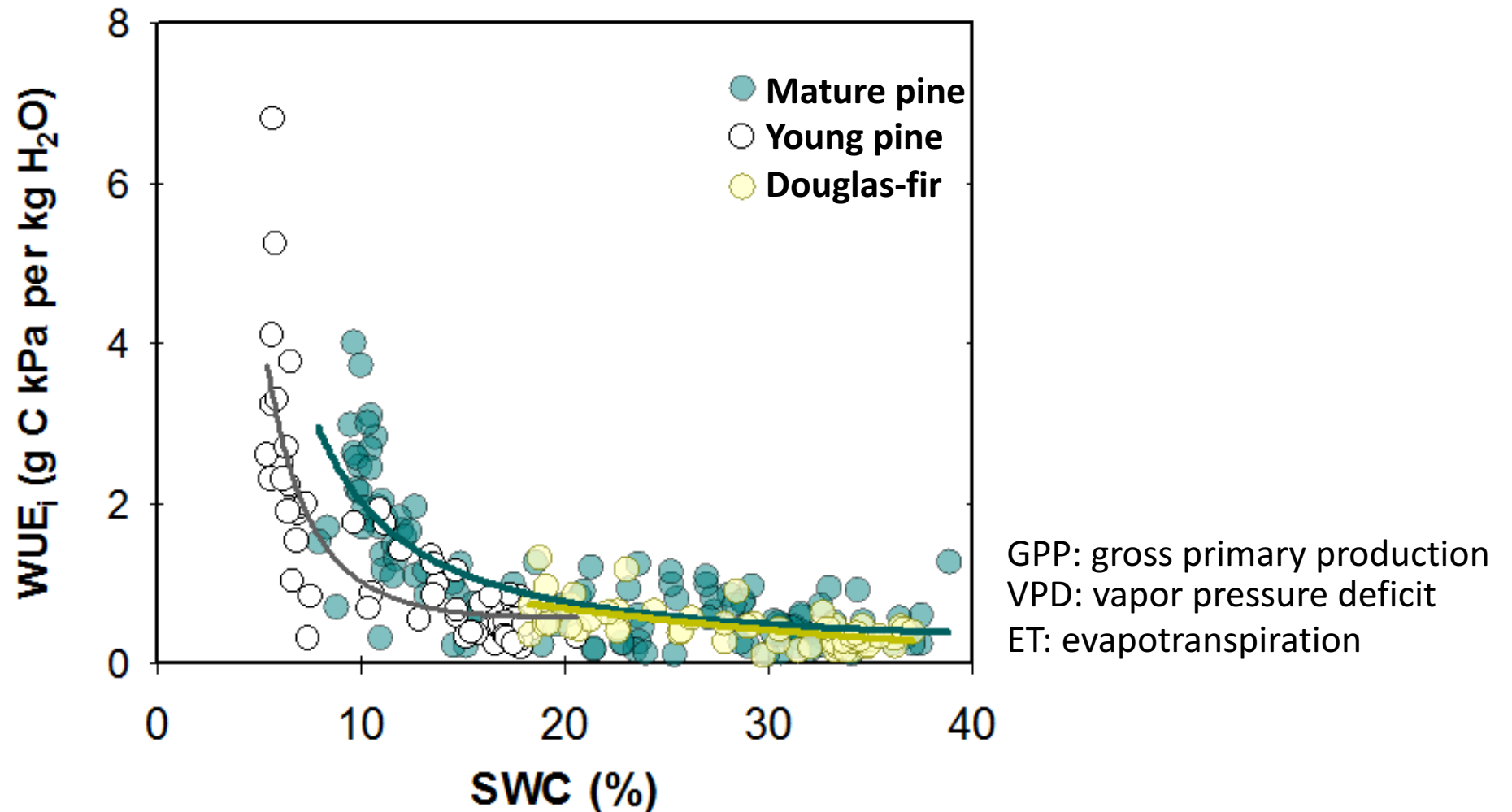
Controlling Mechanism



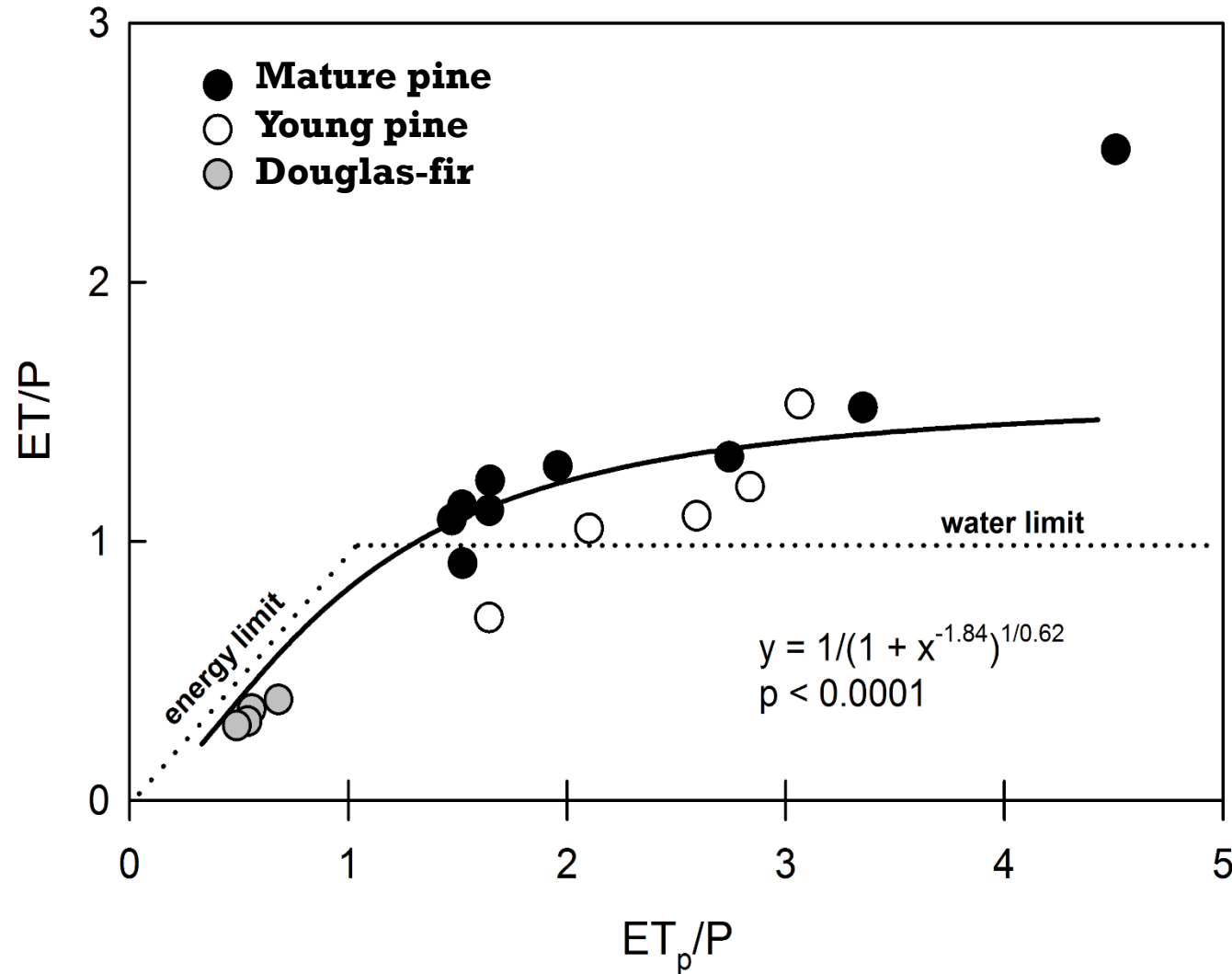
VPD: vapor pressure deficit

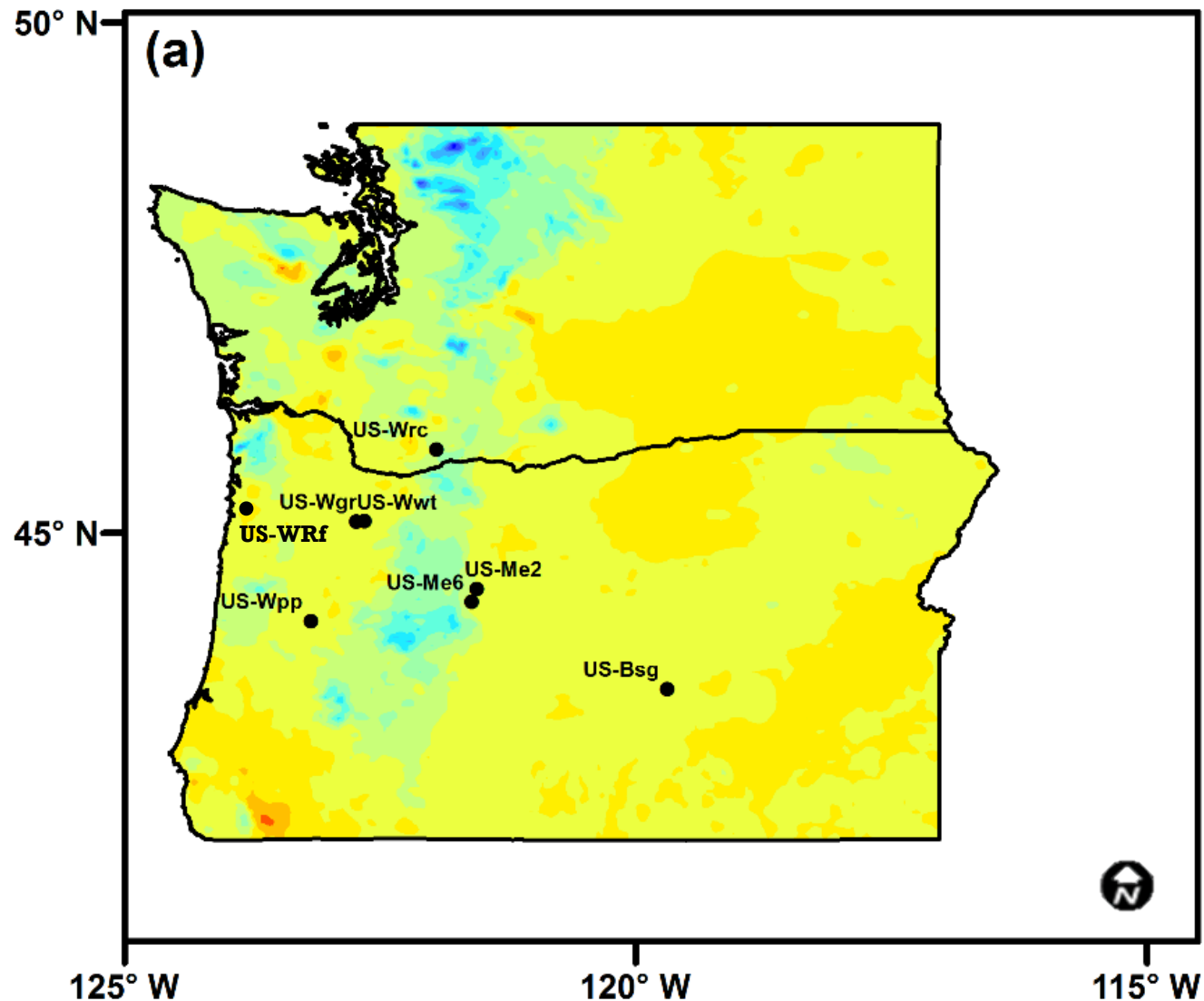


Inherent Water Use Efficiency, $WUE_i = GPP \cdot VPD / ET$

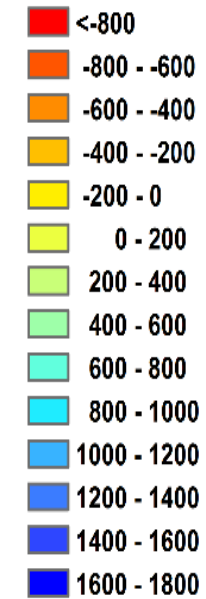


Water Vs. Energy-limited Conditions





Precipitation
anomalies [mm]



US-MRf: Douglas-fir

US-Wgr: grass

US-Wwt: wheat

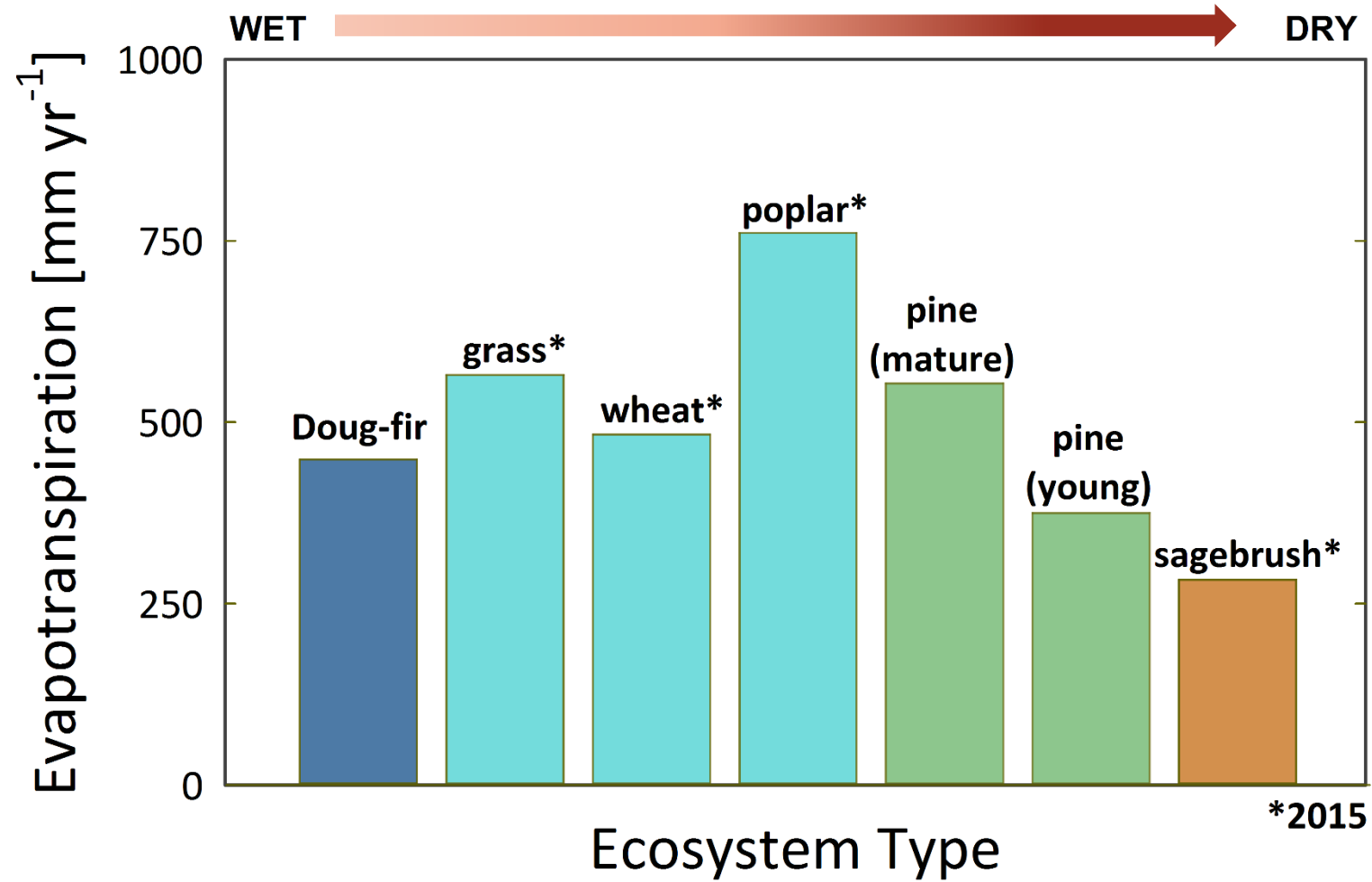
US-Wpp: poplar

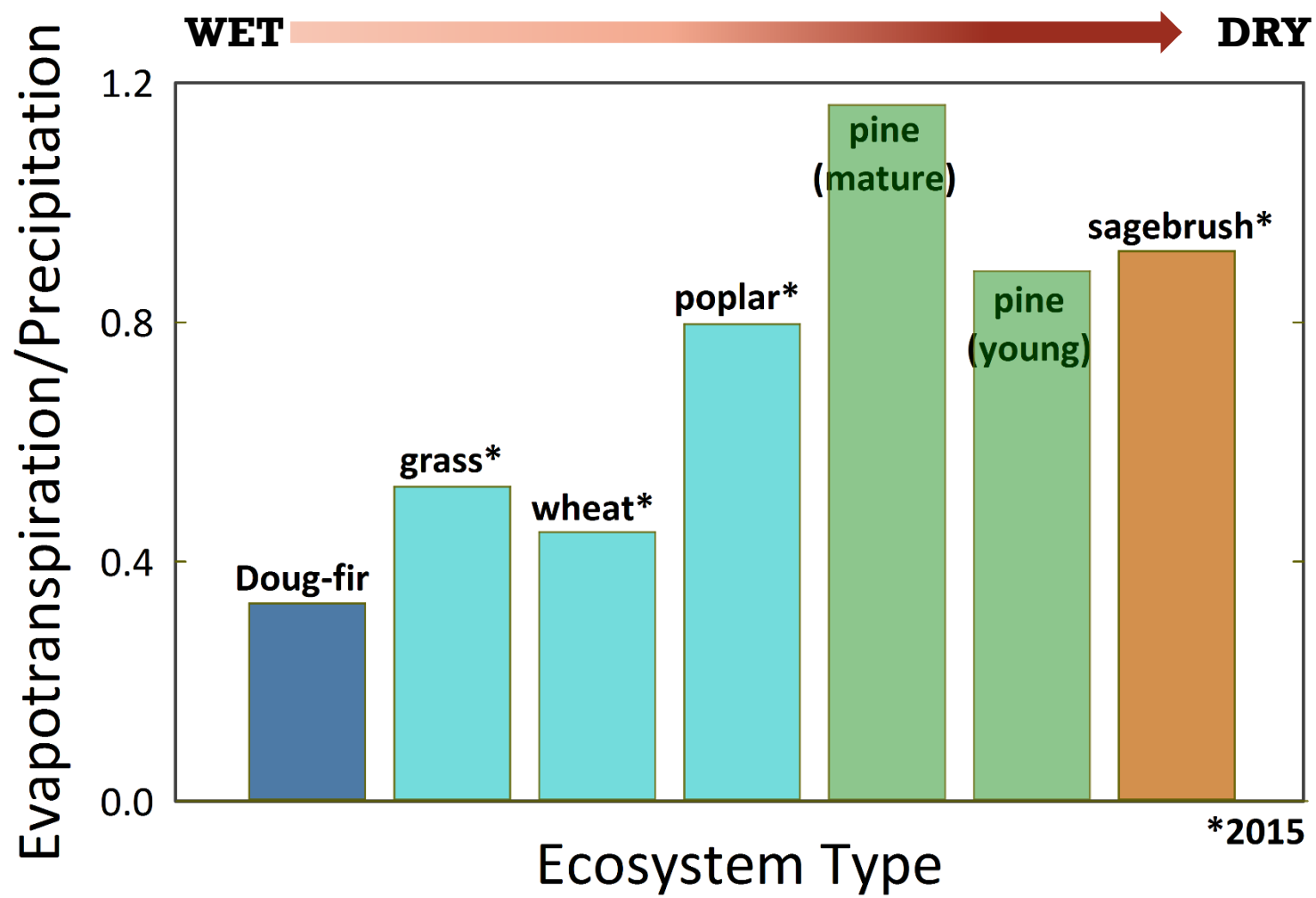
US-Me2: mature ponderosa pine

US-Me6: young ponderosa pine

US-Bsg: sagebrush







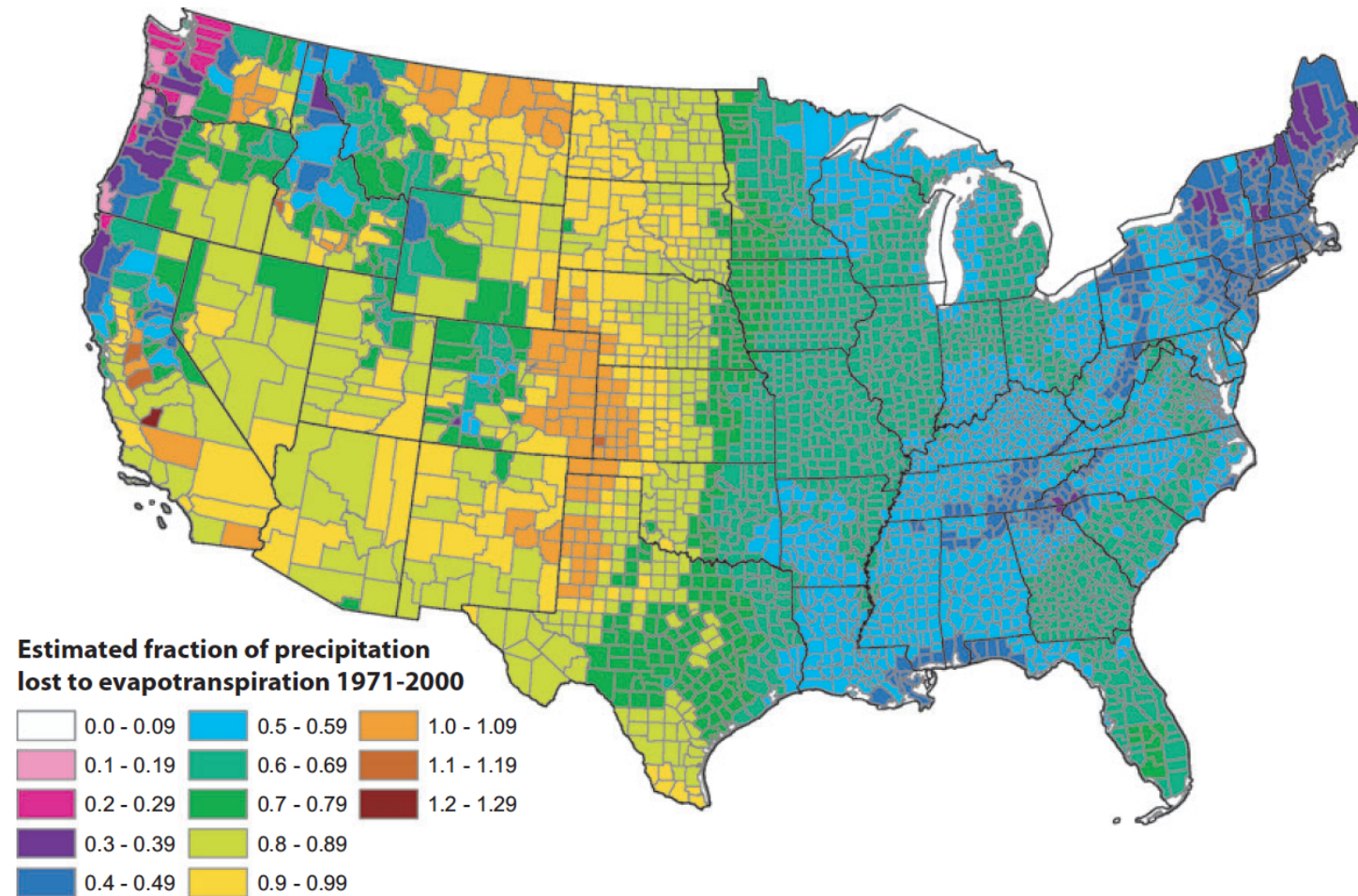


FIGURE 13. Estimated Mean Annual Ratio of Actual Evapotranspiration (ET) to Precipitation (P) for the Conterminous U.S. for the Period 1971-2000. Estimates are based on the regression equation in Table 1 that includes land cover. Calculations of ET/P were made first at the 800-m resolution of the PRISM climate data. The mean values for the counties (shown) were then calculated by averaging the 800-m values within each county. Areas with fractions >1 are agricultural counties that either import surface water or mine deep groundwater.

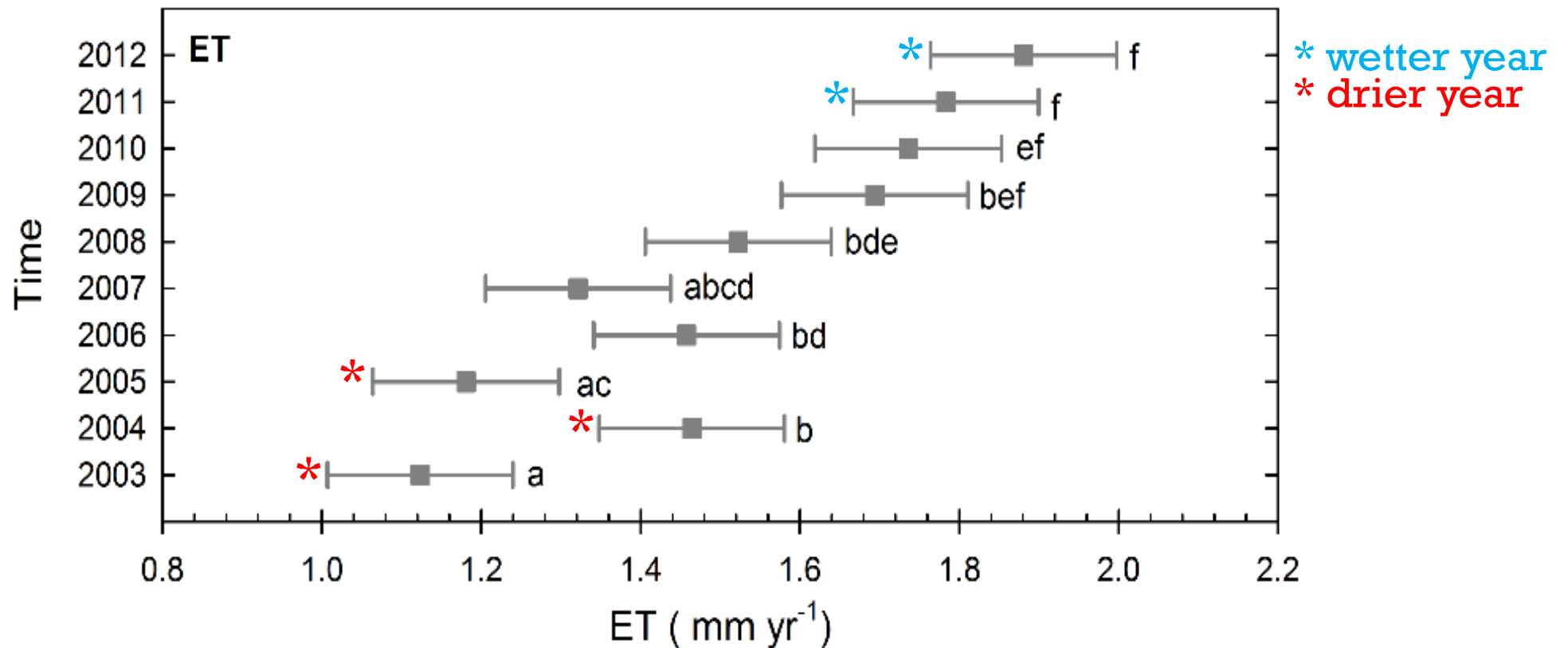


Conclusion

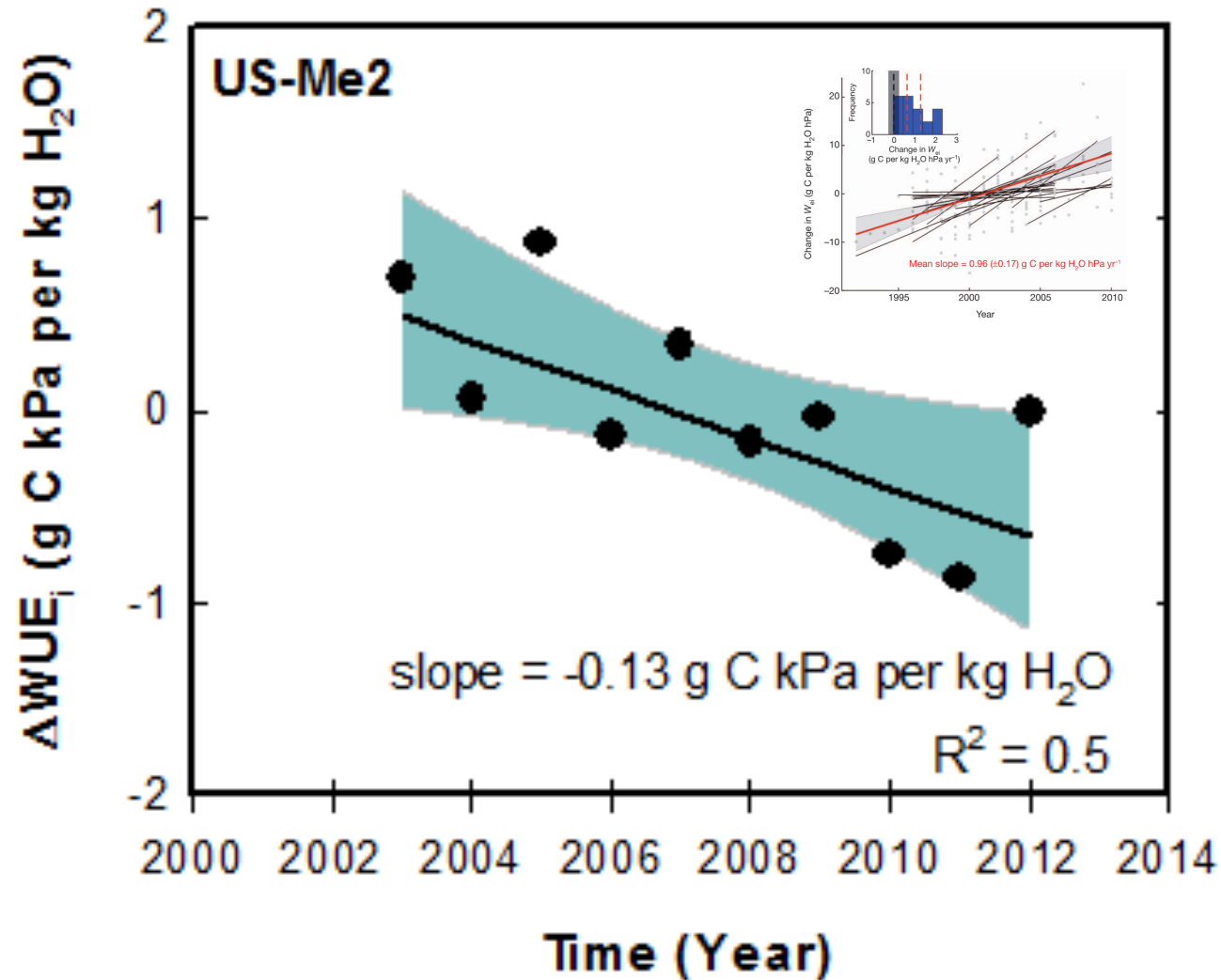
- Individuals of different ages can have markedly different trends in evapotranspiration
- Douglas-fir and ponderosa pine are different enough that sub-plant functional type classes, such as genera, would be a reasonable classification for land system modeling
- Influence of climate and hydrological gradient on evapotranspiration across different land covers is non-linear and often counter-intuitive.



Long-term Variation Of ET



Annual Anomaly Of WUE_i



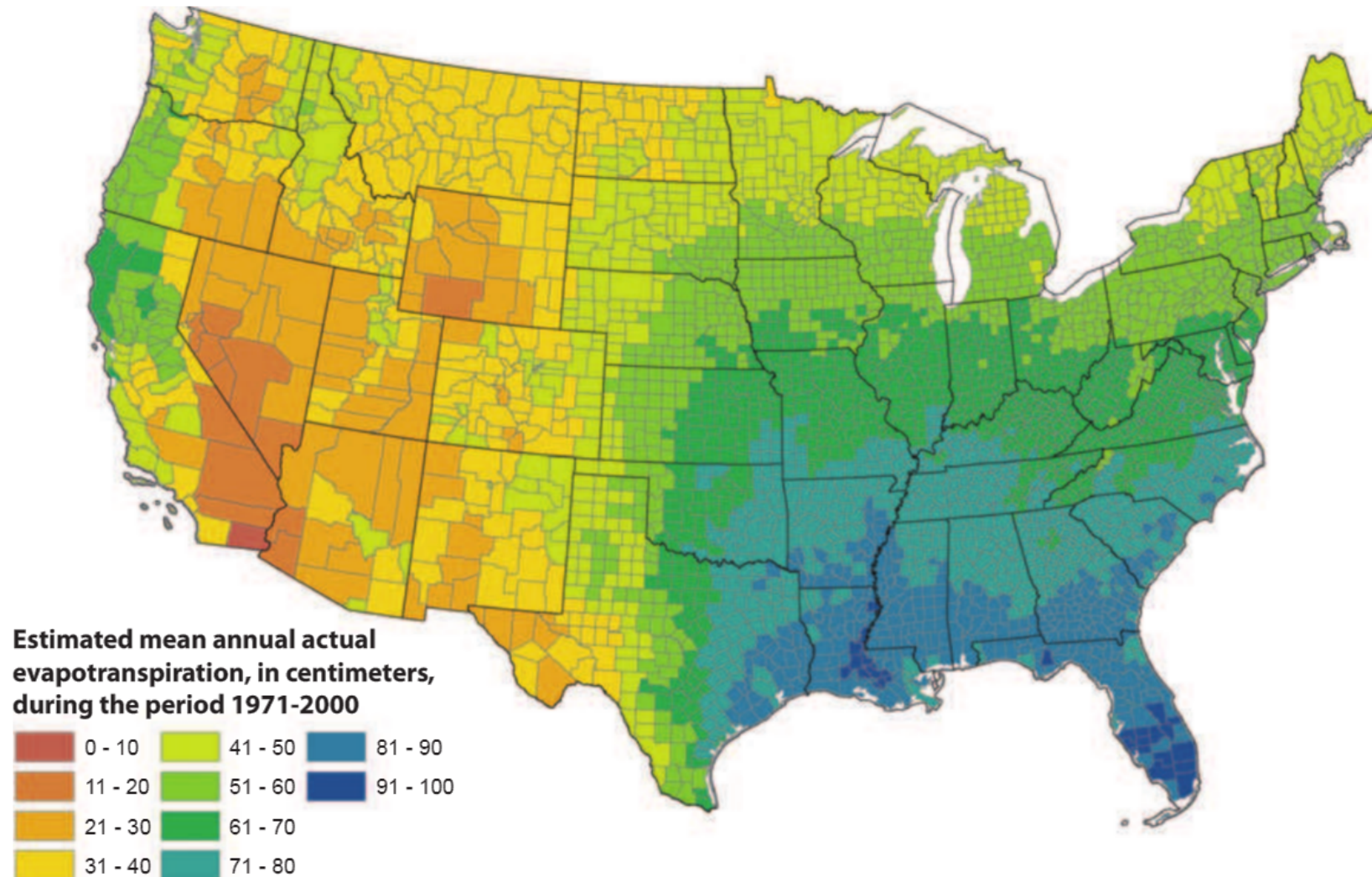


FIGURE 14. Map of Estimated Mean Annual Actual Evapotranspiration (ET) for the Conterminous U.S. for the Period 1971-2000. Estimates are based on the regression equation of ET/P in Table 1 that includes land cover multiplied by the mean annual precipitation from the PRISM climate data for the same period. Calculations of ET were made first at the 800-m resolution of the PRISM climate data. The mean values for the counties (shown) were then calculated by averaging the 800-m values within each county.

